Biostratigraphical and sedimentological aspects of the Cambro-Ordovician strata at the Angosto de Chucalezna: new palynological data (Jujuy, NW Argentina)

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Palabras clave: Estratigrafía. Bioestratigrafía. Cambro-Ordovícico. Palinomorfos. Chucalezna. Jujuy. NW Argentina.

Key words: Stratigraphy. Biostratigraphy. Cambrian-Ordovician. Palynomorphs. Chucalezna. Jujuy. NW Argentina.

Introduction

Cambro-Ordovician strata of the Cordillera Oriental of Jujuy province display widely distributed fossiliferous outcrops, studied by many authors over one hundred years. The Angosto de Chucalezna is one of those localities, with two easy to reach sections with abundant and well preserved fossils, particularly trilobites, palynomorphs and ichnofossils: 1 - The railway section in the western flank of the Rio Grande, and 2- The Quebrada de Moya section eastwards, on the slope of the homonymous peak (Fig. 1).

Paleontological studies have provided several diagnostic elements to date these strata as Upper Cambrian – early Upper Tremadocian (Moya, 1988; Aceñolaza, 1996; Aceñolaza and Albanesi, 1996; Aceñolaza and Poiré, 1998; 1999; Di Cunzolo *et al.*, 2003).

A thick and faulted succession of sandstones and shales of the Mesón and Santa Victoria groups will be examined at this locality. The Cambrian Mesón Group will not be analyzed in detail, but the sedimentologic and paleontologic elements of the Casa Colorada, Alfarcito and Rupasca formations (Santa Victoria Group) will be focused in this opportunity. Tidal and wave dominated successions within a shoreface setting characterize all sequences of this locality, that has been proposed as an ichnological interest point due to the nicely preserved trace fossils associations and the vicinity to a main route (Aceñolaza, 2001a).

In this contribution will be revisited the main aspects of the stratigraphy and biostratigraphy of this locality will be revisited, presenting for the first time a rich and varied palynological association within the upper levels of the Casa Colorada Formation, here referred to the lowermost Tremadocian,

close to the Cambro-Ordovician boundary.

Stratigraphy and Biostratigraphy

The Casa Colorada Formation (Harrington, 1957; emmend. López and Nullo, 1969) display a thick succession of green and black shales and subordinated sandstones. The unit display a lower sandier and an upper shalier interval. The lower sandier sector of the formation has frequently been considered as a different unit (eg. "Alfarcito Formation"), but in this oportunity the entire sequence is understood as one lithological element, as originally defined by Harrington (1957). Following sedimentary characters and paleontological evidence, the Casa Colorada Formation is interpreted as shoreface sedimented strata, within an upper offshore / middle shoreface setting. Deposit feeders are the dominant group reflected by trace fossils, with a proportional less participation of suspension feeders (Di Cunzolo *et al.*, 2003). Sandstone intercalations within the shalier sector of the Casa Colorada Formation and equivalent units were interpreted as storm layers inhabited by oportunistic faunas (Aceñolaza, 1996; Aceñolaza, 2001a, 2001b; Di Cunzolo et al., 2003). Trilobites of the Parabolina (N.) frequens argentina biozone were also recognized in the locality (Aceñolaza, 1996), while monotaxic (brachiopods) and bitaxic (brachiopodtrilobite) bioclasts accumulations have also been described (Aceñolaza and Nieva, 2001).

The Casa Colorada Formation is transitionally followed by the scarcely fossiliferous sandstones and quartzites of the Alfarcito Formation (Harrington, 1957; *emmend*. López and Nullo, 1969). The transitional levels between both units are characterized by an heterolithic sequence of shales and sandstones that have provided a new association of palynomorphs (see data below).

Large to medium scale cross bedded and sigmoidal sandstone bodies with clear reactivation surfaces, infrequent mud drapes, amalgamated and interbedded cross stratified sandstones with frequent herringbone cross-stratification characterize the Alfarcito Formation. Sedimentary strata and ichnological evidence supports a lower shoreface to upper shoreface sedimentary setting. Trace fossils are not abundant, with a clear dominance of suspension feeders ("*Skolithos* ichnofacies"). Fossils are scarce due to the taphonomic and environmental history of strata. The Cambro-Ordovician *Parabolina (N.) frequens argentina* trilobite biozone is also recorded all across strata. Bitaxic bioclast accumulations were also mentioned and its taphonomy analysed (Aceñolaza and Nieva, 2001).

Above the Alfarcito Formation follows the greenish - grayish shales and sandstones of the Rupasca Formation (Harrington, 1957; *emmend*. López and Nullo, 1969). Trilobites of the *Bienvillia tetragonalis-Shumardia (Conophrys) minutula* and *Notopeltis orthometopa* biozones were described from this unit, with *Shumardia (Conophrys) minutula, Hapalopleura clavata, Asaphellus jujuanus, Leptoplastides mariana* and *Protopeltura mesembria* among the most distinctive elements (Aceñolaza, 1996). Conodonts of the *Paltodus deltifer* biozone (early Upper Tremadocian) were recovered from carbonate cemented coquinite levels of the upper part of the sequence in the railway cut: *Drepanodus arcuatus, Drepanoistodus* aff. *concavus, Paltodus deltifer*, Rossodus *tenuis* and *Teridontus nakamurai* (Aceñolaza and Albanesi, 1996; Albanesi and Ortega, 2002). An upper offshore to transitional setting within a shoreface model was interpreted for the unit (Di Cunzolo *et al.*, 2003). Strata is disrupted upwards by faulting.

Palynomorphs of the Casa Colorada formation

A new diversified microfloristic assemblage with relatively good preservation has been recovered from the upper sector of the Casa Colorada Formation (Fig. 5). This is characterized by the dominant presence of taxa belonging to diacromorphitae (*Acanthodiacrodium*, *Ladogella*) acanthomorphitae (*Vulcanisphaera*) and herkomorphitae (*Cristallinium*, *Cymatiogalea*) forms.



Fig. 1. Location map of locality, Chucalezna, Jujuy province.

The following taxa were identified: Acanthodiacrodium achrassi, A. angustum, A. crassus, A. ubui, A. spp., Baltisphaeridium sp., Buedingiisphaeridium tremadocum, Cristallinium cambriense, Cymatiogalea velifera, C. cristata, C. spp., Eliasum sp., Impluviculus multiangularis, Ladogella rommelaerei, Leiosphaeridia sp., Lophosphaeridiumsp., Mychrystridium sp., cf. Nellia longispinata, Nellia? sp., Priscogalea sp., Polygonium dentatum, P. symbolum, P. sp, Saharidia downie, S. fragilis, Timofeevia sp., Trichosphaeridium annovaelensis, T. sp., Veryhachium sp., Vulcanisphaera africana, V. tuberata, V. turbata, and an indetermined Scolecodont.

Most of these taxa were earlier mentioned for either late Cambrian and early Ordovician strata, with many elements crossing the Cambro-Ordovician boundary. It is interesting to point out the lack of typical and exclusive Cambrian taxa, with the exception of *Eliasum* and *Vulcanisphaera turbata*, whose presence may obey to resedimentation. In addition, Fombella Blanco (1982) records *Eliasum* up into the Tremadocian.

The presence of the Ordovician *Acanthodiacrodium angustum* provides interesting elements to the chronological discussion, because this form is usually recorded associated to the guide fossil *Rhabdinopora flabelliformis*.

The early mentioned elements supports a lowermost Ordovician age, close to the Cambro-Ordovician boundary, for the upper levels of the Casa Colorada Formation.

Palynomorphs of the Casa Colorada Formation were also compared to other Cambro-Ordovician associations. No typical and exclusively Cambrian palynomorph associations have been recorded for Argentina yet. In Salta Province Bultynck and Martin (1982) mentioned the presence of Tremadocian forms, sharing some elements with our microflora. Manca *et al.* (1995) and Rubinstein (1997) analyze a more diversified early to late Tremadocian assemblage from Nazareno region (Salta province), presenting different species to those here described. Recently, Aráoz and Vergel (2001), Aráoz *et al.* (2002) and Aceñolaza *et al.* (2003) describe a new association from the Sierra de Zenta (Jujuy and Salta provinces) assigned to the early Tremadocian. This last association shares some elements with the material here described, but the overall proportion of forms do not display a one to one correspondence.

When compared to other palynological associations (e.g. Newfoundland, Canadá: Parsons and Anderson, 2000; Martin and Dean, 1981, 1988; North Africa: Vecoli, 1999), the microflora resembles the association of Newfoundland, Canadá. An important correspondence is recorded with Martin and Dean 's (1981, 1988) latest Cambrian – Tremadocian A6 association (*Arbusculidium destombesii-Vulcanisphaera capillata* assemblage). 10 elements of the early referred association are found in our microflora, highlighting that 5 of them do not appear in the previous Upper Cambrian A5 association (Martin and Dean, 1981; 1988). In a same manner, our material is similar to Parsons and Anderson 's (2000) RA 10ab *Nellia? longispinata-Nellia sukatschevii* microflora, correlated to the previous A6 association of Martin and Dean (1981; 1988).

Palynomorphs of the Casa Colorada Formation share 9 taxa with the "a" and "b" sub-asociations of Parsons and Anderson (2000), highligthing *A. angustum* as the key form to date the lowermost Tremadocian. These data allows a closer correspondence with the RA 10b microflora. The last mentioned authors compared their RA10b sub-association with several zones recognized for the Cambro-Ordovician boundary (eg. China, Russia) and other belonging to the early

Tremadocian (eg. Europe and Africa). These data when compared to our material supports the proposed earliest Tremadocian age for the palynological association of the Casa Colorada Formation. Parsons and Anderson (2000) believe that Cambro-Ordovician boundary microfloras are mainly composed by survival Cambrian assemblages with very few subordinated Tremadocian elements. This situation changed rapidly with the diversification of new taxa in Tremadocian times, reducing the presence of typical Cambrian forms. The palynological association of the Casa Colorada Formation display few survival Cambrian forms with a first Tremadocian association that lacks the posterior diversification of diacromorphitae, herkomorphitae and polygonomorphitae, as happens in all other early recorded palynological associations of the Cordillera Oriental of Argentina.

After comparing the recently established palinozones of Vecoli (1999) for northern Africa, our microflora does not have a full correspondence with those of the Upper Cambrian (*Timofeevia phosphoritica-Dasydiacrodium caudatum* assemblage zone) or the earliest Tremadocian (*Acanthodiacrodium simplex-Arbusculidium destombesii* assemblage zone). With the first assemblage shares *Ladogella rommelaerei*, a typical component of Cambrian associations, but also recorded in Tremadocian associations and several less important biostratigraphic forms as *C. velifera*, *S. fragilis*, *V. africana* and *A. achrassi*. The abundant herkomorphitae and diversified acanthomorphitae in the second association of Vecoli did not allow a closer correspondence.

Taking in account the mentioned basis, we interpret that this new association from the Quebrada de Moya could be considered slightly older than any earlier mentioned association from Argentina, within the lowermost Tremadocian.

Stop 1- Railway cut by the western flank of the Rio Grande

(early Upper Tremadocian Rupasca Formation)



Fig. 2. Stratigraphical section of the railway cut with distribution of facies and fossils.

Thirty nine meters of pale yellowish and greenish shales and sandstones, with thickenning upwards both strata and grain arrangement, are displayed in the railway cut by the route 9 at the Angosto de Chucalezna. The upper sector of the strata displays few coquina type storm layers that provided conodonts of the early Upper Tremadocian *Paltodus deltifer* biozone (Aceñolaza and Albanesi, 1996; Albanesi and Ortega, 2002) (Fig. 2).

The sequence is highly fossiliferous (dominated by trace fossils), and has been several times analysed (Aceñolaza, 1996; Aceñolaza and Albanesi, 1996; Aceñolaza and Poiré, 1999; Di Cunzolo *et al.*, 2003). The infrequent Hapalopleuriid trilobites of Harrington and Leanza (1957) can here be seen (if found).

An heterolithic sequence of mudstones and wave rippled sandstones characterize the lower and middle interval of the section, while the upper sector is distinguished by amalgamated hummocky cross bedded sandstones.

The section display a strong relation between sedimentary facies and trace fossils. The heterolithic facies show a clear dominance of *Teichichnus rectus* and *Trichophycus* isp., while *Cruziana* and *Rusophycus* (*Cruziana furcifera, Cruziana semiplicata*, *Rusophycus polonica*, *Rusophycus latus*, *Rusophycus* isp.) appear in soles of the cross bedded sandstones. *Arenicolites, Didymaulichnus, Dimorphichnus, Phycodes palmatus, Skolithos* and *Monocraterion* are more frequently found in the sandier interval. The last mentioned ichnoassiciation crosscut *Cruziana* and *Rusophycus* (Aceñolaza and Poiré, 1996).

The section is disrupted in the upper and lower sectors by faulting.

Stop 2- Quebrada de Moya section

(Cambro-Ordovician Casa Colorada and Alfarcito formations)

10 km south of Huacalera locality by the eastern flank of the Rio Grande the Quebrada de Moya section display a fairly good faulted sedimentary succession (Fig. 3). The 150 meters thick sequence is represented by the Casa Colorada, Alfarcito and partially the Rupasca formations. Strata are characterized by several intrusives crosscutting its siliciclastic stratigraphy. The Cambrian Mesón Group underlies the entire sequence, while the Rupasca Formation shales and sandstones are tectonically in contact with the overlying Quaternary strata of the Uquía Formation (Aceñolaza, 1996).

Olenid trilobites are dominant fossils in the sedimentary succession. The *Parabolina* (N.) frequens argentina biozone is represented in the Casa Colorada and the scarcely

fossiliferous Alfarcito formations, while the *Bienvillia tetragonalis-Shumardia (Conophrys)* minutula biozone is displayed in the upper Rupasca Formation.

The newly presented palynomorph association was recovered from the upper levels of the Casa Colorada Formation at this section (Fig. 5).

Coquinite levels of the Alfarcito Formation yielded few elements assigned to *Drepanodus arcuatus* (conodonta), but no diagnostic conodonts were found in association (Aceñolaza, 1996).

The fragmentary outcrop of the Rupasca Formation in the Quebrada de Moya belongs to lowermost sector of the unit, while the outcrops of the railway cut by the western flank of the river are placed above the first. Asaphacean and Hapalopleuriid trilobites as well as small lingulids are found in these strata.



1) Drepanoistodus n. sp. A. Sa element, lateral view. Sample CH, CORD-MP 10059, x70.

- Drepanoistodus n. sp. A. M element, lateral view. Sample CH, CORD-MP 10060, x70.
 Rossodus tenuis (Miller, 1980). b element, posterior view. Sample CH, CORD-MP 10061,
- x100.
- Paltodus deltifer pristinus (Viira, 1970). Sa element, lateral view. Sample CH, CORD-MP 10062, x100.
- Paltodus deltifer pristinus (Viira, 1970). M element, lateral view. Sample CH, CORD-MP 10063, x100.
- Drepanodus arcuatus Pander, 1856. f element, lateral view. Sample CH, CORD-MP 10064, x120.
- Teridontus nakamurai (Nogami, 1967). a element, lateral view. Sample CH, CORD-MP 10065, x100.
- Semiacontiodus sp. a clement, postero-lateral view. Sample CH, CORD-MP 10066, x130.
 Utahconus n. sp. A. f element, inner-later view. Sample CH, CORD-MP 10067, x90.
 Paltodus cf. subaequalis (Pander, 1856). Pa element, outer lateral view, Sample CH,
- CORD-MP 10068, x 130.

south-central Sweden, and global correlatives.

All figured specimens are SEM photographs.

Conodont species from the upper part of the Rupasca Formation at Chucalezna section, Cordillera Oriental, Jujuy Province (Albanesi and Aceñolaza, in progress). The conodont association represents the *Paltodus deltifer pristinus* Subzone (*Paltodus deltifer* Zone) of the scheme proposed by Löfgren (1997) for the upper Tremadocian at Brattefors,



Fig. 3. Schematic stratigraphical section of outcrops at the Quebrada de Moya., Chucalezna.



Fig. 5- Palynomorphs of the Casa Colorada Formation at the Quebrada de Moya, Chucalezna. 1. Acanthodiaerodiun angustum (Downie) Combaz 2-Acanthodiaerodium ubui Martin 3-Balthisphaeridium sp. 4-Buedingiisphaeridium tremadoen. Rasul 5-Cristallinium cambriense (Slavikova) Vanguestaine 6-Cymatiogalea velifera (Downie) Martin 7-Cymatiogalea cristale (Downie) Deunff, Gorka and Rauscher 8-Cimatyogalea sp. 9-Eliasum sp. 10-Impluviculus multiangularis (Umnova Volkova 11-Ladogella rommelaerei (Martin) Di Millia, Ribecai and Tongiorgi 12-Ladogella sp. 13- cf. Nellia longispinate Parsons and Andersons, 14-Priscogalea sp. 15-Polygonium dentatum (Timofeev ex Konzalova-Mazancova) Albani 16-Polygonium symbolum Rasul 17-Sabaridia fragilis (Downie) Combaz 18-Scolecodont indet. 19-Timofeevia sp. 20-Trichosphaeridium annoruelensis Timofeev 21-Trichosphaeridium sp. 22-Veryhachium sp. 23-Vulcanisphaera africana (Deunff) Rasul 24-Vulcanisphaera tuberata (Downie 1958) Eisenack et al., 25-Vulcanisphaera turbata Martin (1-8; 11-16; 19-25: 500x - 9; 400x - 10: 1000x - 17: 200x - 18: 100x).

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